Optimal Ranking Regime Analysis of U.S. Summer Temperature and Degree-Days: 1895–2015

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ABSTRACT

The optimal ranking regime (ORR) method was applied to mean summer maximum (TMX_S) and mean summer minimum (TMN_S) temperature and to cumulative summer cooling degree-days (CDD_S) calculated from U.S. climate-division data during 1895–2015. CDD_S is proposed as a proxy for growing degree-days for summer corn given their high rank correlation in station data during 1950–2014. The TMX_S and CDD_S ORR analyses show similar climate-regime patterns. Western and northeastern divisions experienced multidecadal cool periods before 1930 and warm periods after 1990. The 1930s drought appears as decadal warm regimes over the Midwest and Great Plains. Multidecadal TMX_S and CDD_S temperature cycles are evident over the Southeast, but TMX_S and CDD_S variation over the Midwest's Corn Belt agricultural region has been regime free since the early 1940s. By contrast, TMN_S regimes consistent with centennial-scale warming trends are found over most divisions outside the Southeast. From the multidecadal regime patterns detected by the ORR analyses, the TMX_S, TMN_S, and CDD_S series of each climate division were tested for significant linear trends during 1910–2015 and 1970–2015. Significant positive TMN_S trends during 1910–2015 are found in 48 of the 102 divisions, with some western trend magnitudes being greater than 15% of the twentieth-century climatological mean. During 1970–2015, positive TMX_S trends are detected over 39 western and northeastern divisions, but warming TMN_S trends are evident nationally. In some cooler western divisions, positive 1970–2015 CDD_S trend magnitudes exceed 90% of the climatological mean. Consistent with the ORR analyses, Corn Belt TMX_S and CDD_S trends are insignificant during 1970–2015.

1. Introduction

Summer growing season climate plays an important role in U.S. agricultural production. During the summer months, rice is grown in the Mississippi River valley, the Texas Gulf Coast, and the Sacramento River valley. Upland cotton is grown in west Texas, Arizona, the Mississippi River and San Joaquin valleys, and areas of the South. The main sorghum production region includes areas of Kansas, eastern Colorado, and the southern Great Plains. The most concentrated and extensive U.S. corn and soybean production occurs over a midwestern area known as the Corn Belt (USDA 1994, 2017a). During the 2015 and 2016 growing seasons, the United States accounted for

35.5% and 34.1% of global corn and soybean production, respectively (USDA 2017b), most of which was grown in this region. Remote sensing of sun-induced chlorophyll fluorescence indicates that the U.S. Corn Belt during the northern summer is one of Earth's most photosynthetically active areas (Guanter et al. 2014).

Extremes in minimum and maximum temperatures, particularly maximum temperatures, can be a key limiting factor in U.S. summer crop development. Based on regression models using county-level yield and temperature records during 1950–2005, Schlenker and Roberts (2009) found that U.S. corn, soybean, and cotton yields gradually increase as crops are exposed to increasing temperatures below the 29°–32°C range, but that exposure to temperatures above that range resulted in more abruptly decreased yield. Robertson et al. (2013) applied the Schlenker and Roberts (2009) approach to Canadian yield and temperature data and found decreased yield for a range of crops when exposed to temperatures between 28° and 34°C. Ye et al. (2017) found that cumulative exposure to temperatures above

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30°C decreased Kansas corn yields during 1981–2013, with a negative effect on rain-fed yields that is 2–3 times the effect on irrigated yields. Lobell et al.'s (2011) analysis of historical African corn yield and temperature records found that exposure to temperatures above 30°C reduced optimally rain-fed yields by 1% and drought-stressed yields by 1.7%. A summary of temperature threshold yield effects for a variety of crops can be found in Luo (2011).

Although extreme temperatures during important growth phases can hinder plant development and reproduction (Wheeler et al. 2000), crop production has been shown to be linearly related to the accumulated exposure to temperatures between a base temperature and a higher temperature optimal for growth. This has led to the calculation of a derived temperature variable referred to as growing degree-days (GDD), which measures the total exposure to temperatures within that range during a growing season (e.g., Mederski et al. 1973; Peng et al. 1989; Liang et al. 1991).

Given the negative effects of high temperatures on crop production, evidence of long-term U.S. warming trends (Karl et al. 2009; Kunkel et al. 2013) leads to questions about those trends' effects on summer temperature and GDD accumulation. Ideally, work exploring these questions would be based on quality-controlled daily station data from dense meteorological networks providing continuous and long-term data records. Previous research into variation in U.S. temperature extremes has relied on daily data that have been subjected to quality-control (QC) procedures on a case-by-case basis (DeGaetano 1996; DeGaetano and Allen 2002; Peterson et al. 2008; Lee et al. 2014). Such measures are necessary over the United States, because artificial temperature shifts can be introduced by urbanization effects (Karl et al. 1988; Karl and Jones 1989), time-of-observation bias (Karl et al. 1986), and widespread shifts in instrumentation (Quayle et al. 1991). The highest-quality long-term daily dataset provided by the U.S. National Centers for Environmental Information (NCEI) is from the U.S. Historical Climatology Network (USHCN; Menne et al. 2015). The USHCN daily data are derived from a subset of U.S. cooperative station records that have undergone a series of QC checks to detect outliers, artificial recording streaks and gaps, and spatial and temporal inconsistency. However, these data have not been subjected to procedures that can identify and correct for artificial shifts in a station's temperature record relative to those of nearby stations. There is currently no daily dataset provided by the NCEI that is adjusted to correct for non-climatic biases in long-term U.S. temperature records.

While bias-adjusted daily temperature records would be the best data resource to identify long-term variation in U.S. summer extreme temperatures and GDD, the

approach here is to use the improved and bias-corrected monthly data provided by NOAA's Climate Divisional Database (nClimDiv; Vose et al. 2014). Although defined at monthly resolution, nClimDiv data are based on bias-corrected station data and provide continuous records for 344 continental U.S. climate divisions, beginning in 1895. Monthly data are appropriate here, given the emphasis on a ranking analysis of U.S. seasonal temperature since that year. Earlier work based on applying the optimal ranking regime (ORR) method to a previous version of divisional temperature data was limited to seasonally averaged mean monthly temperatures (Mauget and Cordero 2014). The release of nClimDiv data allows for similar ORR analyses here of mean summer maximum temperature (TMX_S) and mean summer minimum temperature (TMN_S). Although this work's focus is on summer temperature variation important to agriculture, nClimDiv data do not include monthly GDD records. The dataset does provide records of monthly cooling degree-days (CDD), which is a functionally similar temperature variable that has fairly high rank correlations with GDD accumulated over the summer months. Summer CDD is considered as a summer GDD proxy here, but the ORR and trend analyses of divisional CDD records may be considered in their own right given the variable's relationship with energy demand (Sailor and Pavlova 2003).

The goal in the following is to explore variation in U.S. summer average maximum and minimum temperature and total summer CDD (CDD_S) and summer GDD (GDD_S) during 1895–2015 using the ORR time series analysis method. Section 2 describes the nClimDiv divisional climate data and the daily USHCN station data used to estimate the rank correlation of CDD_S and GDD_S during 1950–2014. Section 3 describes and demonstrates the ORR method, and section 4 presents the 1895–2015 ORR analyses of climate-division TMX_S and TMN_S . Section 5a presents the GDD_S versus CDD_S rank correlations for 121 USHCN stations, and section 5b presents the 1895–2015 ORR analysis for divisional CDD_S . Section 6 presents the results from linear trend analyses of divisional TMX_S , TMN_S , and CDD_S series during 1910–2015 and 1970–2015. Section 7 presents a summary and discussion.

2. Data

The nClimDiv monthly temperature and CDD values for continental U.S. climate divisions are derived from Global Historical Climatology Network-Daily station temperature data. This process begins by generating a gridded data product from station data with a climatologically aided interpolation process. The resulting gridded data are then

FIG. 1. Locations of the 102 U.S. climate divisions. Red, blue, yellow, and green divisions mark the southeastern, northeastern and southern Florida, midwestern, and western regions, respectively.

area weighted to produce average monthly temperatures over each climate division during 1895–2015. Relative to the previous U.S. climate-division dataset ([Guttman and Quayle 1996](#)), nClimDiv data are based on a station network with improved coverage and quality control. The underlying station data have been adjusted for time of observation bias and were also subjected to bias adjustments ([Menne and Williams 2009](#)) to correct for undocumented changes in observation time, station location and siting conditions, and temperature instrumentation. Monthly CDD data for each climate division were derived by the NCEI from divisional temperature values with the [Thom \(1954, 1966\)](#) algorithm, which estimates the monthly accumulation of daily average temperature above 18.3°C (65°F). Because the 344 NCEI climate divisions have an irregular spatial distribution, with a greater number of divisions defined in the eastern United States relative to the West, monthly mean minimum temperature (TMN), mean maximum temperature (TMX), and CDD values were area-weighted averaged here to the more uniform 102 “megadivision” resolution used by the U.S. Climate Prediction Center ([Fig. 1](#)). TMX_5 and TMN_5 values were calculated by averaging monthly mega-division temperatures over the June–August (JJA) period, and CDD_5 was calculated by summing monthly CDD values over the JJA period.

Daily minimum and maximum temperature station data were obtained from the USHCN. As noted above, although USHCN daily data have been subjected to

QC checks, they have not been bias adjusted for non-climatic temperature shifts. Evaluation here of Corn Belt USHCN daily station data during the summer months shows that a reduction in daily temperature range resulting from the 1980s shift to the Maximum–Minimum Temperature System ([Quayle et al. 1991](#)) is particularly evident. However, the use of these data in [section 5](#) is limited to calculating rank correlations between 121 station’s CDD_5 and GDD_5 totals during 1950–2014. As a result, the USHCN daily data are not used as records of unbiased temperature variation but are used to estimate the rank similarity between CDD_5 and GDD_5 values calculated from the same temperature records.

3. The ORR method

Detecting variation in long-term climate records can be confounded by a time series analysis method’s basic assumptions. Regional climate variability can occur in intermittent regimes or irregular cycles, but indiscriminate linear trend analysis assumes somewhat linear behavior over the periods that trends are fitted. Centennial and multidecadal global temperature records, for example, show evidence of breakpoints and piecewise linear variation with changing trend values [see [Seidel and Lanzante \(2004\)](#) and references therein]. Moving-window-based methods such as wavelet analysis ([Lau and Weng 1995](#)) provide a way of exhaustively sampling a time series’ variation and have been

used to resolve intermittent climate cycles (Wang and Wang 1996; Minobe et al. 2002). However, wavelet analysis and related methods (Harris 1978; Mihovilovic and Bracewell 1991) still assume that a time series' underlying behavior is basically cyclic. By contrast, ORR analysis is based on a less restrictive condition, that is, that climate variability occurs as distinct decadal to multidecadal climate regimes with arbitrary onset and duration. The method detects climate regimes as the most significant sequences of low and high rankings in a time series relative to a null hypothesis that assumes the absence of trends and low-frequency variation (H_1). Its strengths stem from a ranking- and moving-window-based approach that avoids the restrictive assumptions on which some time series analysis methods are based.

In short, the ORR method (Mauget and Cordero 2014; Mauget 2003, 2011, 2015) is a brute-force analysis of rankings to find the most significant sequences of low or high rankings in a time series. Given a 120-yr time series of TMN_S, TMX_S, or CDD_S values, the method begins by ranking a time series' values and then sampling the rankings over running time windows. Each sample of rankings are converted into corresponding Mann–Whitney U and Z statistics (Mann and Whitney 1947), which are normally distributed variables that are proportional to the incidence of high rankings in the sample. Thus, U statistics in a distribution's upper (lower) tail indicate a sample with a high incidence of high (low) rankings. This running sampling of rankings is repeated over a range of sampling sizes to test for regimes of varying length. Finally, the significant Z statistics from all the running analyses are then evaluated to identify the most significant concentrations of low or high rankings, that is, optimal ranking regimes, occurring over nonoverlapping time windows.

Although the U statistic for a sample of rankings can be calculated on the basis of the sample's size and rank sum (Wilks 2011), the statistic is also equal to the total number of data values outside the sampling window that precede each sample value when all values are arranged by rank (Hollander and Wolfe 1999). Thus for a 120-year time series divided into a sample window of $n_I = 20$ years and $n_{II} = 100$ years outside that window, the largest U statistic occurs when the sample contains the 20 highest-ranked years. In that case, 100 nonsample years precede each sample year when arranged by rank, resulting in $U_{\max} = 100 \times 20 = 2000$. The smallest U statistic is produced by a sample with the 20 lowest-ranked years, where no sample ranking is preceded by nonsample rankings when rank ordered ($U_{\min} = 0 \times 20$). Sets of 20 randomly sampled rankings produce U statistics that are normally distributed between those two

extreme values. That distribution's mean is equal to the average of the $U_{\min}$ and $U_{\max}$ values:

$$\mu_0 = 0.5[(0 \times 20) + (100 \times 20)] = 0.5n_I n_{II}. \quad (1)$$

The U distribution's standard deviation can be estimated with (Wilks 2011)

$$\sigma_0 = \left\{ \frac{n_I n_{II} (n_I + n_{II} - 1)}{12} - \gamma \left[\frac{n_I n_{II}}{12(n_I + n_{II})(n_I + n_{II} - 1)} \right] \right\}^{1/2}, \quad (2)$$

where

$$\gamma = \sum_{j=1}^J (t_j^3 - t)$$

is a factor accounting for the incidence of ties in a time series' values, with J indicating the number of groups with tied values and t_j giving the number of tied values within each group. A sample's U statistic can be Z normalized using the null parameters of Eqs. (1) and (2), with significant negative (positive) Z values indicating a significant incidence of low (high) rankings relative to a null hypothesis that assumes random sampling (H_0):

$$Z_0 = \frac{U - \mu_0}{\sigma_0}. \quad (3)$$

Although the ORR method tests for significant low-frequency climate variability, the Monte Carlo–derived null parameters that specifically test the H_1 hypothesis in previous work (Mauget 2003, 2004) can result in Z statistics that are less conservative, that is, with larger magnitudes, than those resulting from Eq. (3). To prevent assigning significance to ranking sequences consistent with random variation, the approach here was to compare the Eq. (3) Z_0 magnitude with that of a second Z_1 statistic that specifically tests the H_1 hypothesis:

$$Z_1 = \frac{U - \mu_1}{\sigma_1}. \quad (4)$$

For each time series and each n_I sample size, the Eq. (4) μ_1 and σ_1 null parameters were calculated through a Monte Carlo process that repeatedly sampled rankings from 120-yr autoregressive noise series consistent with H_1 . The Monte Carlo procedure used to generate H_1 -consistent U null distributions and calculate their μ_1 and σ_1 null parameters is described in Mauget and Cordero (2014). As in that analysis, a ranking sample's Z statistic was assigned on the basis of which Z normalization

resulted in the smallest Z magnitude and the most conservative significance level:

$$|Z| = \min(|Z_0|, |Z_1|). \quad (5)$$

Figures 2a–e demonstrate the ORR procedure applied to Fig. 1's division 84 (northern California) TMX_S time series. Figure 2a shows the TMX_S series and its low-frequency variation derived through a low-pass Lanczos filter (Duchon 1979) with a cutoff frequency of $\nu = (20 \text{ yr})^{-1}$. Figure 2b shows the series of Eq. (5) Z statistics for TMX_S rankings sampled over running 20-yr periods, with horizontal lines marking the Z statistics' negative and positive significance at two-sided 95%, 99%, and 99.9% confidence levels and Z magnitudes > 4.0 . Figure 2b also defines the shading scheme used to identify those eight Z levels on the figure's left axis. Figure 2c's colored horizontal regime bars mark the duration of the 20-yr ranking regimes found to be negatively or positively significant at a 95% or better confidence level in Fig. 2b, superimposed on the TMX_S series. Those bars are vertically placed according to each ranking regime's Z statistic, which is measured on the figure's right axis, and are shaded with the Fig. 2b color scheme.

The ORR method identifies optimally significant and nonoverlapping sequences of data rankings by calculating U and Z statistics from rankings sampled over moving time windows. Those windows range here from 10 to 40 years in width. For each of the 31 window sizes, running series of Z statistics like that of Fig. 2b are calculated with Eq. (5) for that window size for each time series. For each running Z series, the corresponding probability p statistics are then subjected to Ventura et al.'s (2004) false-detection-rate procedure. This test's demonstration can be found in the appendix of Mauget (2018). In addition to testing significance of U statistics calculated with a fixed sample size, Z normalization also makes it possible to order the significance of U statistics calculated using different sample sizes. After Z normalizing the running U statistics from each analysis, the positive and negative Z statistics from all 31 Z series that pass the Ventura et al. (2004) p -threshold test are combined as in Fig. 2d. Within those pooled statistics, the ranking sequences that produce the highest absolute Z values over nonoverlapping time periods are identified (Fig. 2e). For more details of the screening process that identifies a time series' optimally significant regimes, which is illustrated in Figs. 2d and 2e, see section 3 of Mauget and Cordero (2014).

By detecting what might be considered a basic feature of climate variation—here, summer temperature and CDD regime periods—the ORR method can identify a wide range of climate variability. In Fig. 2e, the northern

California TMX_S series shows an irregular multidecadal temperature cycle between 1900 and the mid-1950s, which is found in both the Lanczos low-pass TMX_S trace and alternating cold (1897–1916), warm (1926–40), and cold (1941–56) ORR periods. After the mid-1950s summer maximum temperatures are relatively stationary until the late 1990s but then increase into a warm ORR period during 2000–15. Figure 3 shows additional examples of how the method can detect arbitrary climate variation. In Fig. 3a, highly significant ($|Z| > 4.0$) periods of low- and high-ranked temperatures during the early (1900–39) and late (1977–2015) years of the southern Florida (division 29) TMN_S record mark a long-term warming trend. As shown in Fig. 3b's western Colorado (division 61) TMN_S series, however, widely separated cold and warm regime periods may not always indicate a trend between those periods. Figure 3c's Southern California (division 88) TMN_S series shows an example of a “wavy trend,” that is, a multidecadal cycle before the 1960s that is superimposed on a longer-term warming trend. A warming trend is also detected in Fig. 3d's northern Alabama (division 18) TMN_S series, but over a multidecadal period between a cool 1955–76 regime and a warm 1993–2015 regime.

4. ORR analyses of TMN_S and TMX_S

The ORR method identifies climate regimes occurring over nonoverlapping time windows. As a result, a time series' most significant regimes can be plotted on a horizontal line referred to as a Z line, with the regime's positive or negative significance marked by the Fig. 2b color scheme. Vertically condensing a time series's regime periods this way makes it possible to plot ORR results for numerous time series on a single graph and to identify common patterns of decadal to multidecadal variability.

Figure 4 shows the TMX_S Z lines for each of Fig. 1's 102 climate divisions during 1895–2015. Each line's vertical coordinate is assigned according to the division's number on the Fig. 1 map, which also divides the continental United States into four areas. Divisions 1–27 are located over Fig. 1's red southeastern region. Divisions 28–36 define the blue region, which includes seven northeastern divisions and two divisions in southern Florida. Divisions 37–59 define a yellow region that coincides with the major agricultural production areas of the midwestern and northern Great Plains states. Figure 4's green-shaded Z lines show TMX_S regimes for divisions 60–102, which are located in the western United States, the Rocky Mountain region, and parts of the western Great Plains.

In Fig. 4, significant concentrations of low-ranked TMX_S conditions before 1930 are detected over western (Fig. 4a) and northeastern (Fig. 4b) divisions, and

FIG. 2. (a) Summer (JJA) TMX_s record for northern California (division 84) during 1895–2015. The heavy trace shows low-frequency variation derived by using a Lanczos low-pass filter with a $\nu = (20 \text{ yr})^{-1}$ cutoff frequency. (b) Mann–Whitney Z statistics of ranked TMX_s values sampled over running 20-yr time windows. The horizontal lines and left-axis shading scheme show two-sided 95% ($Z = \pm 1.96$), 99% ($Z = \pm 2.575$), and 99.9% ($Z = \pm 3.29$) confidence levels and $|Z| > 4.0$. (c) As in (a), but with the horizontal width of the colored bars showing the 20-yr ranking regimes in (b) with $|Z| > 1.96$. The bar's vertical position shows corresponding Z values as marked by the

FIG. 3. As in Fig. 2e, but for (a) southern Florida (division 29) TMN_s , (b) western Colorado (division 61) TMX_s , (c) Southern California (division 88) TMN_s , and (d) northern Alabama (division 18) TMN_s .

regimes of high-ranked conditions are found after 2000 over most of those same divisions (Figs. 4c,d). In Fig. 4e, warm summers associated with the 1930s drought appear as highly significant runs of high-ranked TMX_s conditions over midwestern climate divisions 46–60, and climate divisions in the western Great Plains and interior West (divisions 63–74, 80–84). During the 1940s and 1950s, the final cool phase of Fig. 2e's pre-1960

multidecadal cycle can be seen over neighboring climate divisions 68–94 (Fig. 4f). Although summers of the 1930s were not significantly high-ranked in many of these divisions (e.g., divisions 77–80), this suggests a similar multidecadal TMX_s cycle over the western United States before 1960. Southeastern divisions 1–27 show evidence of a multidecadal TMX_s cycle, with warm regimes before the late 1950s shifting to cooler summer

←

right axis. (d) As in (c), but with significant TMX_s ranking regimes from ORR analyses conducted with sampling windows ranging from 10 to 40 yr. (e) The optimally significant sequences of low-ranked and high-ranked TMX_s occurring over nonoverlapping time windows.

FIG. 4. The Z lines for the ORR of each climate division's TMX_S record during 1895–2015. The Z line's vertical coordinates are assigned according to the Fig. 1 division numbers. Colored regions on both vertical axes mark lines for Fig. 1's four shaded regions. Positive and negative significance at 95%, 99%, 99.9% confidence levels and $|Z| > 4.0$ are marked by the Z -scale shading scheme defined at the top. The black frames labeled with a–h outline TMX_S regime features that are discussed in the text.

regimes during the early 1960s to mid-1990s (Fig. 4g). The latter cool regime interrupts the general twentieth-century U.S. warming trend and as a result has been referred to a “warming hole” (Pan et al. 2004; Portmann et al. 2009; Meehl et al. 2012). In addition to the Southeast's distinctive warming-hole variation, that region shows little evidence of significant regimes of high-ranked TMX_S values after the late 1990s like those found in western

and northeastern climate divisions (Figs. 4c,d). Similarly, in contrast to western, northeastern, and southeastern regime variation, over the Midwest there are no clear patterns of either high-ranked or low-ranked TMX_S regimes between the early 1940s and 2015 (Fig. 4h).

Figure 5 is Fig. 4's counterpart for mean summer minimum temperatures. When compared with Fig. 4, regime patterns consistent with long-term positive TMN_S trends

FIG. 5. As in Fig. 4, but for the ORR analyses of each division's TMN_5 rankings during 1895–2015. The black frames labeled with a–f outline TMN_5 regime features that are discussed in the text.

are more apparent over most of the continental United States outside the Southeast. These trends are marked by multidecadal and highly significant ($|Z| > 4.0$) regimes of low-ranked TMN_5 in the early twentieth century (Fig. 5a) and similarly significant regimes of high rankings after the 1980s (Fig. 5b). Over western climate divisions 68–97, decadal cool regimes are outlined in Fig. 5c during the 1940s and 1950s. Given these division's previous cool conditions before 1930, and warm conditions after 1980, this long-term variation is consistent with the wavy trend

shown in Fig. 3c's division 88 time series. Centennial TMN_5 trend patterns are also found in some midwestern divisions, for example, divisions 36–39 and 43–46, where consistent TMX_5 regime patterns in Fig. 4 are limited mainly to the warm 1930s period. Similar to Fig. 4e, midwestern warm regimes are also detected in the TMN_5 rankings during the 1930s (Fig. 5d), but occur at generally lower significance levels and appear over a more limited area of the Great Plains (divisions 50–60, 66–69). However, southeastern climate divisions that show almost

FIG. 6. (a) Spearman rank correlation values for summer (JJA) cooling degree and corn growing degree totals derived from 121 USHCN daily station temperatures during 1950–2014. (b) Histogram of the r_s values mapped in (a).

no evidence of recent warm TMX_5 conditions in Fig. 4 show a consistent pattern of warm TMN_5 regimes after the mid-1990s (Fig. 5e). Relative to Fig. 4h, Fig. 5f's warming-hole southeastern cool regimes are more consistently significant, with low-ranked TMN_5 conditions more concentrated from the mid-1950s to mid-1970s.

5. Summer degree-day analysis

Both CDD and GDD values measure the cumulative exceedance of daily average temperature above a specified base temperature T_b over the course of daily, monthly, or seasonal time periods. The former is used to estimate electrical demand for air conditioning (e.g., Downton et al. 1988; Sailor and Pavlova 2003), and the latter estimates cumulative exposure to temperature conditions suitable for crop development (Luo 2011). The following demonstrates that the rank correlations

of CDD₅ and GDD₅ values calculated from daily USHCN TMX and TMN values may be high enough that, for the purposes of the ORR method, the NCEI divisional CDD data might be used as a proxy for 1895–2015 GDD variation.

a. GDD and CDD rank correlation

To calculate CDD₅ and GDD₅ totals from the same daily temperature records, 121 stations were selected from the 1218-station USHCN network that had no more than 14 missing daily TMX or TMN values in any of the JJA periods of 1950–2014 (Fig. 6a). These stations also had no more than 1% missing JJA data over all those 65 summer seasons. Missing daily minimum and maximum temperatures were filled in with stochastically generated values (Hanson et al. 1994).

Following Lobell et al. (2013) and Ye et al. (2017), daily growing and cooling degree-days were calculated

from interpolated hourly temperature values T_h that were estimated by fitting 24-h cosine functions to daily TMN and TMX values:

$$T_h = \frac{\text{TMX} + \text{TMN}}{2} + \frac{\text{TMX} - \text{TMN}}{2} \times \cos[\pi(h_{\max} - h)/12], \quad h = 1, 24. \quad (6)$$

Maximum daily temperatures were assumed to occur 2 h before sunset ($h_{\max}$), and a day's integer sunset hour at each station location was estimated with procedures provided in Allen et al. (1998).

As described above, cooling degrees measure the accumulation of temperature above a base temperature threshold. The CDD threshold generally coincides with the temperature at which regional electrical demand is minimized, which varies geographically (Ranson et al. 2014). This threshold ($T_{b\text{-cdd}}$) was set to 18.3°C (65.0°F) in keeping with NCEI CDD calculation and previous studies (Downton et al. 1988; Sailor and Muñoz 1997; Sailor and Pavlova 2003). To calculate daily CDD values, hourly temperatures in each day's interpolated T_h record below 18.3°C were first set to the 18.3°C T_b floor value:

$$T_{h\text{-cdd}} = \max(T_h, 18.3). \quad (7)$$

A day's hourly CDD values (CDD_h) are the resulting adjusted $T_{h\text{-cdd}}$ values minus the 18.3°C threshold:

$$\text{CDD}_h = T_{h\text{-cdd}} - T_{b\text{-cdd}}. \quad (8)$$

CDD values for each day of the year (doy) during JJA are the average of the day's 24 CDD_h values:

$$\text{CDD}_{\text{doy}} = \frac{1}{24} \sum_{h=1}^{24} \text{CDD}_h, \quad (9)$$

and CDD_S , with units of accumulated degrees Celsius above the threshold temperature (DD_{18}), is the sum of the 1 June–31 August daily CDD values:

$$\text{CDD}_S = \sum_{\text{doy}=152}^{243} \text{CDD}_{\text{doy}}. \quad (10)$$

Total GDD_S were similarly calculated for each station during 1950–2014, but with each day's hourly temperatures subjected to an additional ceiling temperature requirement. Daily growing degrees for crops are accumulated between two temperature thresholds that vary by crop (Luo 2011). Here, GDD values for each day of the summer growing season (GDD_{doy}) were calculated using threshold values for corn, which is the most widely produced U.S. feed grain. The corn $T_{b\text{-gdd}}$ base threshold temperature, that is, the minimum temperature required

for crop development, was set to the 8°C (46.4°F) value assumed in Butler and Huybers (2013) and Ye et al. (2017). The corn optimum ceiling temperature, that is, the temperature above which crop development and yields are negatively impacted, was set to the 29°C (84.2°F) value estimated by Schlenker and Roberts (2009). Thus for each day's 24-h interpolated T_h record, temperatures below 8°C were set to that $T_{b\text{-gdd}}$ floor value, and values above 29°C were set to a ceiling of 29°C. A day's hourly corn GDD values GDD_h are these adjusted T_h values ($T_{h\text{-gdd}}$) minus the 8°C base threshold:

$$\text{GDD}_h = T_{h\text{-gdd}} - T_{b\text{-gdd}}. \quad (11)$$

Similar to Eqs. (9) and (10), GDD_{doy} values for each day of the year during JJA are the average of the day's 24 GDD_h values, and total GDD_S is the sum of the 1 June–31 August GDD_{doy} values.

Although daily CDD_{doy} and GDD_{doy} values are calculated with different base temperatures, and CDD_{doy} calculation has no ceiling temperature threshold, the CDD_S and GDD_S records for most of Fig. 6a's stations have fairly high rank correlations during 1950–2014. To show the agreement between both variables' rank ordering, Fig. 6a maps the Spearman rank correlation r_s calculated from each station's 65 CDD_S and GDD_S rankings during that time, and Fig. 6b is a histogram showing the r_s distribution. Of the 121 r_s values, 111 (92%) are greater than 0.90, with 85 values (70%) in the 0.92–0.96 range. Of the 10 r_s values that are less than 0.90, 9 are greater than 0.89. Given these generally high rank correlations, ORR analyses that are based on CDD_S rankings derived from available NCEI divisional data should be qualitatively similar to analyses of GDD_S . The use of CDD_S as a GDD_S proxy here does not require these variables to be quantitatively similar but does require ordinal similarity. ORR analyses conducted on the GDD_S and CDD_S rankings of Fig. 6a's stations on the basis of running 10–25-yr windows during 1950–2014 (Figs. S1, S2, and S3 in the online supplemental material) show that this is generally the case. As a result, the Fig. 7 analysis of CDD_S derived from divisional data may also be broadly representative of GDD_S variation during 1895–2015.

b. ORR analyses of CDD_S

Figure 7 plots the Z lines for each climate division's CDD_S rankings. When compared with the ORR analyses of TMX_S and TMN_S in Figs. 4 and 5, the patterns of CDD_S variability closely follow those of TMX_S . As in Fig. 4, low-ranked CDD_S conditions are detected over western (Fig. 7a) and northeastern (Fig. 7b) climate divisions before 1930. High-ranked conditions are also

FIG. 7. As in Fig. 4, but for the ORR analyses of each division's CDD_5 rankings during 1895–2014. The black frames labeled with a–h outline CDD_5 regime features that are discussed in the text.

found after 1980 over most of those same divisions (Figs. 7c,d). The warm summers of the 1930s produced highly significant concentrations of both TMX_5 (Fig. 4e) and CDD_5 (Fig. 7e) over the northern Great Plains (divisions 51–59). Over the West, alternating periods of cool conditions before 1930 (Fig. 7a), warm conditions over divisions 81–84 in northern California and Nevada during the 1930s (Fig. 7e), and cool regimes during the 1940s and 1950s (Fig. 7f) are similar to the cyclic TMX_5 regime patterns of Figs. 4a, 4e, and 4f. However, over

Fig. 7c's western climate divisions, the onset years of recent warm CDD_5 periods tend to precede those of the warm TMX_5 regimes in Fig. 4c. In Fig. 7e, sequences of high-ranked CDD_5 mark the warm summers of the 1930s, but with effects more concentrated in the Great Plains (divisions 44–60 and 62–71). The southeastern multidecadal temperature cycle in Fig. 4f is also evident in CDD_5 in Fig. 7f, but with more consistent effects of the warming-hole period during the 1960s and 1970s. Last, as in Fig. 4g, there is a conspicuous lack of

significant CDD_S ranking regimes over the midwestern climate divisions between 2015 and the 1930s drought period (Fig. 7g).

6. Linear trend analyses for 1910–2015 and 1970–2015

In Figs. 5b and 5e, significant warm TMN_S regimes are almost uniformly found over the continental United States after 1980. In some divisions in the West (e.g., divisions 70–94), southern Florida (divisions 28 and 29), the Northeast (e.g., divisions 30–36), and the Midwest (divisions 37–39, 43–46), these regimes are multidecadal and highly significant ($Z > 4.0$). Before 1930, many of those same divisions show extended and similarly significant cool regimes that are roughly centered on the year 1910 in Fig. 5a, which can indicate centennial-scale warming trends (e.g., Fig. 3a,c). But over southeastern climate divisions, the cold and warm TMN_S regimes of Figs. 5f and 5e suggest a more recent warming trend since the 1970s. In contrast to the recent warmth in minimum summer temperature over almost all of the United States, warm TMX_S and CDD_S ORR periods after 1980 are limited to Fig. 1's western (Figs. 4c and 7c) and northeastern climate divisions (Figs. 4d and 7d). Over southeastern divisions 1–27, there is relatively little evidence of either TMX_S or CDD_S warm regimes after 1990 in Figs. 4 and 7. To show the location and magnitudes of possibly significant centennial summer warming trends, Figs. 8a–c map the normalized linear trends of TMN_S, TMX_S, and CDD_S during 1910–2015. To estimate the magnitudes of more current trends, Figs. 8d–f map similarly normalized trends for 1970–2015. The following description of the significance test for linear trends assumes a 1910–2015 trend analysis.

Trend significance was determined following Ebisuzaki's (1997) Monte Carlo–based method. This approach estimates the significance of linear trends in four steps:

- 1) Detrend the 1910–2015 time series and calculate its Fourier transform.
- 2) Using the resulting Fourier amplitudes and a randomized phase value, generate a noise series with the same duration, mean, and variance as the detrended data series by using the inverse transform and calculate its linear trend.
- 3) Repeat the second step 10 000 times to calculate 10 000 null trend values consistent with trendless variation.
- 4) Given the mean and standard deviation of the step-3 null statistics, normalize the undetrended series 1910–2015 linear trend into corresponding Z and p statistics.

The p values resulting from each of the 102 climate division's trends were subjected to the p -ranking procedure

of Ventura et al. (2004), which limits the false discovery rate (FDR) in the 102 tests to no greater than 5%. The test limits the FDR, that is, the percentage of significant trend values that are falsely rejected, by imposing lower and more restrictive p thresholds to determine significance. For an FDR of 5%, these p thresholds are less than 0.05. For a more in-depth explanation and demonstration of this test, see Ventura et al. (2004) or Wilks (2016).

For the TMX_S, TMN_S, and CDD_S trends during 1910–2015, the p -ranking procedure set p thresholds of 0.012, 0.022, and 0.016, respectively. The same thresholds for the 1970–2015 trend analyses were 0.018, 0.038, and 0.032. To normalize significant trend magnitudes from divisions with widely different summer temperature climates, trend magnitudes were divided by the variable's twentieth-century climatological mean C_{20C} . A trend's 1910–2015 magnitude is measured as the difference between a significant linear trend line's 2015 (T_{2015}) and 1910 (T_{1910}) values. Thus the normalized 1910–2015 trends τ_{1910} are those magnitudes, divided by the variable's 1901–2000 mean C_{20C} :

$$\tau_{1910} = 100.0 \frac{T_{2015} - T_{1910}}{C_{20C}}. \quad (12)$$

Likewise, normalized 1970–2015 trends τ_{1970} are the 1970–2015 trend magnitude divided by the twentieth-century mean:

$$\tau_{1970} = 100.0 \frac{T_{2015} - T_{1970}}{C_{20C}}. \quad (13)$$

Figures 8a–c map the divisional τ_{1910} values for TMX_S, TMN_S, and CDD_S, and Figs. 8d–f map those variables' τ_{1970} values. Figures 8g–i map each variable's twentieth-century mean.

In Fig. 8a, positive τ_{1910} values for TMX_S range from 2.5% (division 28) to 6.5% (division 94), with significant trends found in the Southwest and interior West, Florida, and the northeastern seaboard. Significant negative τ_{1910} trends between −4.8% and −3.7% are found in five midwestern divisions. These are the only significant negative trends found in all three variables during either 1910–2015 or 1970–2015. Of 102 climate divisions, only 24 show significant TMX_S trends of either sign. In Fig. 8b, 48 of 102 climate divisions show significant positive TMN_S trends during 1910–2015. Significant TMN_S trends are found in the West and interior West, climate divisions adjacent to the Great Lakes and in the Northeast, and in southern Florida. The largest TMN_S τ_{1910} value (23.0%) is found in western Colorado (division 86). Other TMN_S τ_{1910} values of greater than 17.6% are found in Idaho (division 72), Washington (division 76), Oregon (division 80), and California (division 86). Twenty-five climate

FIG. 8. Normalized 1910–2015 linear trends for (a) JJA TMX, (b) JJA TMN, and (c) cumulative JJA CDD, and normalized 1970–2015 linear trends for (d) JJA TMX, (e) JJA TMN, and (f) cumulative JJA CDD; gray-shaded divisions mark insignificant trends. Also shown is the 1901–2000 climatology of (g) mean JJA maximum temperature, (h) mean JJA minimum temperature, and (i) cumulative JJA CDD.

divisions in the West, Midwest, and Northeast have τ_{1910} trends above 10.7%.

In Fig. 8c, significant τ_{1910} values for CDD_S vary between 13.5% (west Texas division 102) and 189.2% (the Oregon coast division 78). Note that Figs. 8c and 8f's scales are nonlinear, with 10% increments below 90% and 50% increments above 90%. Other western climate divisions with τ_{1910} values above 90% include Washington state divisions 76 and 77 and western Colorado division 61.

These are regions with low C_{20C} values ($<150 \text{ DD}_{18}$) in Fig. 8i, which indicates large CDD_S trends during 1910–2015 relative to low climatological means. Other divisions with proportionately large CDD_S trends ($\tau_{1910} > 70\%$) include divisions in Oregon (division 79), Idaho (division 72), and Wyoming (division 64).

In Fig. 8d, normalized 1970–2015 TMX_S trends vary from 2.3% (southern Florida division 29) to 7.5% (Idaho division 71). When compared with Fig. 8a's pattern of

positive τ_{1910} values for TMX_S , there is more consistent evidence of larger 1970–2015 TMX_S trend magnitudes in Fig. 8d. In the former, there are 7 interior western climate divisions with τ_{1910} values above 4.8%; in the latter, there are 20 divisions between southern Texas and the Pacific Northwest with τ_{1970} above that value. By contrast, Eastern Seaboard and Florida climate divisions with significant positive TMX_S trends after 1970 show τ_{1970} values that are below 4.8%.

In Fig. 8e, consistent with the nearly nationwide shift to warm TMN_S regimes after the early 1980s (Figs. 5b,e), significant positive TMN_S trends are found in 95 of the 102 divisions during 1970–2015. The corresponding τ_{1970} values vary from 3.9% (eastern Iowa division 47) to 15.6% (New Mexico division 99). Similar to the higher τ_{1970} percentages for TMX_S in Fig. 8d, almost all of the largest ($>10.7\%$) values for TMN_S in Fig. 8e are found in climatologically cooler ($TMN_S C_{20C} < 10^\circ\text{C}$) divisions along the Pacific coast, in the Pacific Northwest, and in the interior West. By contrast, Great Plains and mid-western divisions to the east show relatively weak ($<6.1\%$) τ_{1970} values for TMN_S .

In Fig. 8f, significant τ_{1970} values for CDD_S vary from 9.4% (the southern division 14) to 219.7% (the Oregon coast division 78). As in Fig. 8c, the lowest (highest) τ_{1970} values tend to occur over divisions with the highest (lowest) C_{20C} values in Fig. 8i. Thus, over southern climate divisions with twentieth-century mean CDD_S values $> 750 \text{ DD}_{18}$, τ_{1970} values are generally less than 20%. As mean CDD_S decreases in mid-Atlantic, Great Lakes, and northeastern climate divisions, τ_{1970} values increase into the 20%–40% range. This inverse relationship is most clear over climate divisions along the northwestern Pacific coast and high-elevation divisions in the interior West with low climatological CDD_S values. Thus, for example, the twentieth-century climatological mean for division 78 is $13.0 \text{ DD}_{18} (^\circ\text{C})$, while the linear trend's magnitude, that is, $T_{2015} - T_{1970}$, is 28.7°C . During JJA, 2014 and 2015 total CDD_S for that climate division rose to 63.3 and 72.8 DD_{18} , the two highest values for 1895–2015. Over western Colorado division 61 ($\tau_{1970} = 161.4\%$), the C_{20C} mean is 27.6 DD_{18} and the linear trend's rise during 1970–2015 is 44.5°C . The 10 most highly ranked CDD_S values for that climate division during 1895–2015, which varied between 65.6 and 120.6 DD_{18} , occurred after 1993.

7. Summary and discussion

To identify trends and decadal to multidecadal regimes in mean summer (JJA) TMX_S , TMN_S , and cumulative CDD_S , the ORR time series analysis method (section 3) was applied to U.S. climate-division data

during 1895–2015. Given the generally high rank correlation with summer growing degree-day totals for corn estimated from station data during 1950–2014 (section 5a), the CDD_S variation found here in the divisional data is proposed as a proxy for divisional GDD_S variation during 1895–2015.

The ORR analyses of TMX_S (Fig. 4) and CDD_S (Fig. 7) show generally similar variation. Early twentieth-century cool regimes are found over divisions in the West, the Northeast, and Florida (Figs. 4a,b and 7a,b), whereas most of those same divisions experienced warm regimes after 1990 (Figs. 4c,d and 7c,d). These are regions where U.S. divisional summer mean temperatures were significantly correlated with Northern Hemisphere summer temperatures during 1896–2012 according to Mauget and Cordero (2014). Thus this long-term temperature variation appears to be the U.S. expression of broader hemispheric variability. Both TMX_S and CDD_S show a low-frequency multidecadal cycle over the southeastern United States (Figs. 4h and 7h) associated with twentieth-century warming-hole variation (Portmann et al. 2009; Meehl et al. 2012; Mauget and Cordero 2014). Over the Midwest, however, both the TMX_S and CDD_S analyses (Figs. 4g and 7g) show a near absence of significant summer warm or cold regimes since the 1930s drought period (Figs. 4e and 7e).

By contrast, the TMN_S ORR analyses (Fig. 5) show evidence of centennial-scale warming trends in most climate divisions outside the southeastern United States. This consists of highly significant ($|Z| > 4.0$) multidecadal periods of low-ranked TMN_S in the early twentieth century (Fig. 5a) and high-ranked regimes of similar duration and significance after 1980 (Fig. 5b). Over the Southeast, the warming hole's cool regime is apparent during the 1950s, 1960s, and 1970s (Fig. 5f). However, although the TMX_S and CDD_S analyses of Figs. 4 and 5 show little evidence of recent southeastern warmth, warm TMN_S regimes are evident over the Southeast after the early 1980s (Fig. 5e).

Linear trend analyses of the three temperature variables during 1910–2015 and 1970–2015 are generally consistent with the regime patterns that are seen in Figs. 4, 5, and 7. Although significant positive TMX_S trends are found in the West, the Northeast, and Florida during 1910–2015 (Figs. 8a), positive TMN_S trends over that period (Fig. 8b) are more consistently evident in those regions and in Midwest and Great Lakes climate divisions. Relative to twentieth-century climatology (Figs. 8g,h), the magnitudes of 1910–2015 TMN_S trends [Eq. (12)] generally exceed those of TMX_S . In Fig. 8a, the largest TMX_S trend magnitude as a percentage of climatology is 6.5%; in Fig. 8b, the largest TMN_S trend magnitude is 23.0%. Over northwestern and interior West climate

divisions with low (<150 DD₁₈) mean twentieth-century CDD_S values (Fig. 8i), the 1910–2015 CDD_S trend magnitudes exceed 70% of climatology (Fig. 8c). The CDD_S trend magnitudes for climate divisions along the Oregon and Washington coasts are 189.2% and 179.5% of climatology, respectively.

The normalized TMX_S trend magnitudes for 1970–2015 [Eq. (13)] show significant positive trends over the West and Texas, Florida, and Eastern Seaboard climate divisions (Fig. 8d). Relative to 1910–2015 TMX_S trends, positive TMX_S trends since 1970 are more consistently present between southern Texas and the Northwest and have larger magnitudes (4.8%–7.5%). However, 63 of the 102 climate divisions, including midwestern and Great Plains agricultural divisions, show insignificant TMX_S trends after 1970.

Consistent with the nearly national evidence of warm TMN_S regimes after 1980 (Figs. 5b,e), 95 of the 102 climate divisions show significant positive TMN_S trends during 1970–2015 (Fig. 8e). Like the 1970–2015 TMX_S trends, the largest TMN_S magnitudes as a percentage of climatology are found in the interior West and the Northwest. During 1970–2015, significant positive CDD_S trends are found in 79 of 102 climate divisions, with insignificant trends in the northern Great Plains, the Midwest, and northern Maine (Fig. 8f). Like the 1910–2015 CDD_S trends, the more recent CDD_S trend magnitudes in some western divisions are substantial ($>90\%$) relative to climatology.

The summer temperature trends found here are consistent with previous work that has found greater warming in U.S. minimum temperatures relative to maximum temperatures (Karl et al. 1993; Gallo et al. 1999; Lee et al. 2014). The proposed causes for this differential warming include changes in cloudiness, cloud cover height, atmospheric water vapor, aerosol loading, and atmospheric circulation (Karl et al. 1984, 1988; Vose et al. 2005). Warming in minimum temperatures can also be a distinguishing urban heat island (UHI) effect (Hausfather et al. 2013) that can be particularly evident over the summer months (Oke 1982; Karl et al. 1988).

The station data on which nClimDiv divisional data are based have been adjusted for nonclimatic bias with the Menne and Williams (2009) pairwise homogenization (PH) algorithm. Although this algorithm does not explicitly correct for urbanization, Hausfather et al. (2013) demonstrate that it does effectively reduce the difference between U.S. urban and rural temperature trends in monthly USHCN station data. They estimate that, nationally, urban warming effects account for 14%–21% of the increase in unadjusted minimum temperatures during 1895–2010 and 6%–9% during 1960–2010. After PH, they found that the algorithm effectively removes UHI biases

in minimum temperatures after 1960 and reduces biases before 1930 by 50%. These reductions, however, apply to nationally averaged temperature effects that were calculated from all monthly data during 1895–2010. As a result, they may not reflect the PH algorithm's effect on regional or seasonal UHI biases. Thus, uncorrected urbanization effects may contribute to some degree to the centennial-scale TMN_S trends found in Figs. 3a and 3c, Figs. 5a and 5b, and Fig. 8b. Given Hausfather et al.'s (2013) observation that residual UHI signals not removed by the PH algorithm are less significant after 1960, and the relatively weak UHI effects on maximum temperatures (Karl et al. 1988; Hausfather et al. 2013), 1970–2015 TMX_S trends may be less affected. Barring an analysis similar to that of Hausfather et al. (2013) conducted seasonally over regional spatial scales, however, the magnitude of uncorrected UHI effects on the TMN_S and TMX_S trends found here are uncertain.

The highest-magnitude trends as a percentage of climatology were found in 1970–2015 CDD_S trends. From an air-conditioning (AC) market survey of U.S. cities during the mid-1990s, Sailor and Pavlova (2003) derived a function that relates annual CDD to municipal-level per capita AC use. This function's nonlinearity [their Eq. (4)] suggests that in cooler regions with climatologically low CDD, modest CDD increases can lead to disproportionate increases in AC use. Thus, 20%–40% CDD_S trends over densely populated northeastern climate divisions (Fig. 8f) may also foreshadow a developing trend toward greater summer electrical consumption. Even larger ($>80\%$) CDD_S trends over the Pacific Northwest suggest the possibility of similar trends in Portland, Oregon, and Seattle, Washington, consumption.

In Figs. 8d–f, the strongest TMX_S, TMN_S, and CDD_S trends during 1970–2015 are found outside major U.S. agricultural regions, that is, northwestern coastal, interior western, and northeastern climate divisions. Over California's Central Valley (divisions 84 and 87) TMX_S trends are insignificant and significant positive TMN_S trend magnitudes are less than 10% of climatology. However, Central Valley CDD_S trends (26.2% and 26.9%) are more substantial. Recent TMX_S trends for southern climate divisions are weak ($<3.0\%$) or insignificant. Southern TMN_S trends during 1970–2015 are less than 8.4% of climatology, and CDD_S trend magnitudes are under 20%. Great Plains and midwestern TMX_S and CDD_S trends during 1970–2015 are insignificant and positive TMN_S trends are weak relative to other U.S. regions. In addition, the ORR analyses of TMX_S and CDD_S show no clear regime patterns over midwestern climate divisions since the 1930s (Figs. 4g, 7g). Given the high levels of rank correlation between CDD_S and GDD_S in station data during 1950–2014 (Fig. 6), this

suggests that GDD_S over this important U.S. growing region has also been stable during the past seven decades.

Estimating agriculturally relevant climate variation from seasonal data has its limitations. Although seasonal variability was emphasized here, temperature-related crop stress impacts can occur over submonthly time scales (Suwa et al. 2010). Daily minimum temperature records might better resolve the subseasonal effects of positive TMN_S trends evident over almost all of the continental United States in recent decades. Long-term bias-adjusted daily temperature records could be used to confirm GDD_S variation inferred here from CDD_S data, for example, positive GDD_S trends over California's Central Valley. Daily records could also be useful in studying long-term variation in the length of the U.S. summer growing season. A version of nClimDiv divisional data defined at daily resolution that is currently under development at the NCEI (I. Durre 2017, personal communication) may help to explore these issues in future work.

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