

On the Effect of Fluctuations in Vertical Acceleration on the Coalescence Growth of Drizzle Drops

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The general expression for the terminal velocity for a spherical water drop of radius y is

$$U = (2g\rho y^2/9\eta)(24/\text{Re}C_D), \quad (1)$$

where η is the viscosity of air, g the acceleration of gravity, ρ the density of the water drop, Re the Reynolds number, and C_D the drag coefficient. For droplets with radius $< 45 \mu$, $(24/\text{Re}C_D) \approx 1$ and Stoke's law provides a fairly good approximation for the terminal velocity under a variety of environmental conditions as verified by experimental results of Beard and Pruppacher (1969). For droplets with radius $> 45 \mu$ up to a

radius of 1000μ , the inverse of that factor can be expressed as a linear empirical function of the radius (Sulakvelidze *et al.*, 1967), i.e.,

$$\text{Re}C_D/24 = 1.52 \times 10^2 (\rho_p/\rho_0)y, \quad (2)$$

where ρ_0 and ρ_p are the air densities, respectively, at sea level and at the cloud level with pressure p . Consider the case of a drizzle-size drop, say, with initial radius 200μ falling through a cloud represented by an empirical Khrgian and Mazin spectrum

$$n(y) = 1.45 \times 10^{-6} (L/y_{av}^6) y^2 \exp(-3y/y_{av}),$$

with liquid water content L of 1 gm m^{-3} , an average radius y_{av} of 7.5μ , and a droplet size range from 3 to 29μ . The coalescence growth rate of such a drizzle-size drop can be described by

$$dy_l/dt = (\pi/3) \int_0^{y_l} y_s^3 n(y_s) Y_c^2 (U_{yl} - U_{ys}) dy_s, \quad (3)$$

where y_l and y_s are the drizzle-drop and cloud droplet radii, respectively, Y_c the linear collision efficiency as defined by Shafri and Neiburger (1963), and $(U_{yl} - U_{ys})$ the terminal velocity difference between the drizzle drop and smaller cloud droplets. Expressing the Stokes velocity as $U_{ys} = 1.28 \times 10^3 g y_s^2$, and substituting (2) into (1), the drizzle terminal velocity becomes a linear function of y_l and $U_{yl} = 9.45 g y_l$, if we choose a cloud level at 900 mb with virtual temperature 0C . By introducing these expressions for the terminal velocities, (3) can be written as

$$dy_l/dt = g(\pi/3) \int_0^{y_l} n(y_s) (A y_l - B y_s^2) Y_c^2 dy_s, \quad (4)$$

where $A = 9.45 \text{ sec cm}^{-1}$ and $B = 1.28 \times 10^3 \text{ sec cm}^{-2}$.

The coalescence growth history of a typical drizzle drop obtained from integrating (4) numerically as shown in Table 1 for an initial radius of 200μ . This result can be fitted well by an exponential function, with the form of (4) suggesting the structure

$$y = y_0 \exp(gbt), \quad (5)$$

where y and y_0 are the drizzle drop radius at time t and initial time, respectively (we are herewith omitting the subscript l), and b is a parameter determined by the least-square criterion in fitting the growth curve with (5). In general, the parameter b depends on the small droplet spectrum and for the present case $b = 1.9964 \times 10^{-6}$.

Spatial variation of the velocity field inside a cloud is an observed fact. For instance, updrafts and downdrafts are known to be both present in the mature stage of a thunderstorm. Vertical air velocity records from aircraft traverses of developing cumuli (Warner, 1970) also show marked small-scale fluctuations. In such a cloud environment the individual drops are likely to be subjected to the influence of fluctuating vertical accelerations.

A reasonable assumption is that the fluctuating vertical acceleration experienced by the drizzle drop has a

normal distribution with mean zero and standard deviation σ . It will superpose on the acceleration of gravity with the resultant acceleration X being a random variable with the probability density function:

$$f_x(x) = [\sigma(2\pi)^{1/2}]^{-1} \exp[-(x-g)^2/(2\sigma^2)]. \quad (6)$$

The large drop radius now becomes a function of X , i.e.,

$$Y = y_0 \exp(btx) = h(x). \quad (7)$$

Since $y = h(x)$ is differentiable for every x with $h'(x) > 0$ for all x , and since X is a continuous random variable, then $Y = h(X)$ is a continuous random variable with probability density function given by

$$f_Y(y) = f_x[h^{-1}(y)] \left| \frac{d}{dy} h^{-1}(y) \right| = [bt\sigma(2\pi)^{1/2}]^{-1} \times \exp\{-[\log y - bt(C+g)]^2/(2b^2t^2\sigma^2)\}, \quad y \geq y_0$$

$$= 0, \quad \text{otherwise} \quad (8)$$

where $h^{-1}(y) = \log(y/bt) - C$ and $C \equiv (y_0/bt)$.

The expected value of the collector drop radius can be found using the important fact that the expectation of a function of a random variable is equal to the expectation of the function with respect to the random variable; therefore,

$$E(Y) = \int_{-\infty}^{\infty} y_0 \exp(btx) [\sigma(2\pi)^{1/2}]^{-1} \times \exp[-(x-g)^2/(2\sigma^2)] dx = [y_0 \exp(gbt)] \exp[\frac{1}{2}(\sigma bt)^2]. \quad (9)$$

It is seen that the expected radius is that radius the large drop would reach with no fluctuations in acceleration times an amplification factor, $\exp[\frac{1}{2}(\sigma bt)^2]$. For the case of a large drop with the initial size 200μ falling through the aforementioned cloud, if the standard deviation of the fluctuating field is equal to $g/10$, which perhaps represents the upper limit that is realizable in natural clouds, the radius increase will be 0.5% at 600 sec and 1.2% at 800 sec over those predicted by the growth equation. Therefore, the effect of a fluctuating vertical acceleration on the accelerated growth of drizzle drops should be unimportant and negligible under realizable cloud conditions. Fig. 1 shows the logarithm of the probability density function of Y at $t = 600 \text{ sec}$ with initial large drop size and cloud spectrum as mentioned before and with $\sigma = g/10$. For comparison, the drop would grow to 632μ at $t = 600 \text{ sec}$ under the same conditions but without the presence of vertical acceleration fluctuations.

TABLE 1. Growth history of drizzle drop with initial radius 200μ in a cloud having a Khragian-Mazin spectrum with $L = 1 \text{ gm m}^{-3}$, $y_{av} = 7.5 \mu$, and size range $3-29 \mu$.

t (sec)	0	100	200	300	400	500	600	700	800
r (μ)	200	241	291	352	427	519	632	771	942

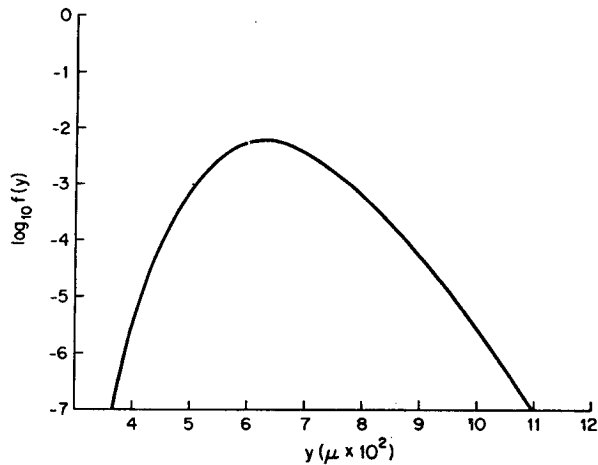


FIG. 1. $\log_{10} f(y)$ vs y at $t=600$ sec for vertical acceleration fluctuations normally distributed with $\sigma=g/10$, $y_0=200 \mu$.

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REFERENCES

- Beard, K. V., and H. R. Pruppacher, 1969. A determination of the terminal velocity and drag of small water drops by means of a wind tunnel. *J. Atmos. Sci.*, **26**, 1066-1072.
- Shafir, U., and M. Neiburger, 1963. Collision efficiencies of two spheres falling in a viscous medium. *J. Geophys. Res.*, **68**, 4141-4148.
- Sulakvelidze, G. K., *et al.*, 1967: *Formation of Precipitation and Modification of the Hail Process*. Israel Program for Scientific Translations, Jerusalem.
- Warner, J., 1970: The microstructure of cumulus cloud. Part III. The nature of the updraft. *J. Atmos. Sci.* **7**, 682-688.