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Supplemental Material

Journal of Hydrometeorology

A Framework for Diagnosing Factors Degrading the Streamflow Performance of a Soil Moisture
Data Assimilation System

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Supplemental Material

29 **A framework for diagnosing factors degrading the streamflow performance of a soil**

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moisture data assimilation system

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36 **S1. Mathematical details of ensemble Kalman filter (EnKF)**

37 The ensemble Kalman filter (EnKF) method is one of the most commonly used data
38 assimilation (DA) techniques in hydrology. EnKF was first introduced by Evensen [1994] and
39 has subsequently been applied to a large number of land DA applications [e.g., Crow and Ryu,
40 2009; Chen et al., 2014; Massari et al., 2015]. It represents model error by an ensemble of model
41 run replicates, which are combined with measurements sequentially to update model states.
42 Specifically, the EnKF method is based on a propagation model and a measurement model:

43
$$x_{k+1} = f(x_k, u_k) + \omega_k \tag{S1}$$

44
$$\tilde{y}_k = Hx_k + v \tag{S2}$$

45 where subscript k is a discrete time index; x is a column vector of model states to update (the
46 column vector length is the total number of state variables to update); u is model meteorological
47 forcing; $f()$ is a land surface model that propagates states to the next timestep; ω lumps together
48 modeling errors during propagation from various sources including forcing data error, model
49 structure error and parameterization error; \tilde{y} is measurement data, in our context surface SM
50 measurements; H is an observation operator that relates model states x to measurements \tilde{y} ; and v
51 is measurement error.

52 In a standard EnKF, an ensemble size of N model replicates is propagated and updated
53 sequentially over time in the following way:

54 1) An ensemble of initial model states is first generated by perturbing the initial deterministic
55 model states to represent initial state error;

56 2) For each ensemble member, the land surface model is run until the next measurement time
 57 with perturbed meteorological forcing to represent forcing error. Model states are directly
 58 perturbed as well to represent random errors from model structure and parameterization;

59 3) Once an observation time is reached, the Kalman gain K is calculated as:

$$60 \quad K_k = P_k H^T \cdot (H P_k H^T + R)^{-1} \quad (\text{S3})$$

61 where R is the measurement error variance, and the forecast state error covariance matrix P_k is
 62 estimated by sampling across the propagated ensemble states:

$$63 \quad P_k = \frac{1}{N-1} \sum_{j=1}^N (\hat{x}_k^{-(j)} - \bar{\hat{x}}_k^-)(\hat{x}_k^{-(j)} - \bar{\hat{x}}_k^-)^T \quad (\text{S4})$$

64 where $\hat{x}_k^{-(j)}$ is the propagated state vector at time k for the j th ensemble member, and $\bar{\hat{x}}_k^-$ is the
 65 mean of $\hat{x}_k^{-(j)}$ across all ensemble members;

66 4) Following the calculation of K , each ensemble member of states is individually updated as:

$$67 \quad \hat{x}_k^{+(j)} = \hat{x}_k^{-(j)} + K_k \cdot (\tilde{y}_k + v_k^{(j)} - \hat{y}_k^{-(j)}) \quad (\text{S5})$$

68 where $\hat{y}_k^{-(j)}$ is the simulated measurement at time k for the j th ensemble member, i.e.,

69 $\hat{y}_k^{-(j)} = H \hat{x}_k^{-(j)}$; $v_k^{(j)}$ is random noise added to represent measurement error whose error statistic
 70 is consistent with R in Equation (S3).

71

72 **S2. Mathematical details of the evaluation metrics used in the study**

73 1) Percent error reduction (PER)

74 PER is defined as the percent reduction in the root-mean-squared error (RMSE)
 75 compared to the open-loop baseline:

$$76 \quad PER = \left[1 - \frac{RMSE_a}{RMSE_{open}} \right] \times 100 \quad (S6)$$

77 where the subscripts *open* and *a* denote the open-loop and DA analysis runs, respectively.

78 2) Kling-Gupta efficiency (KGE)

79 The Kling-Gupta efficiency (KGE) [Gupta et al. 2009] combines the performance in
 80 terms of correlation, variance and bias:

$$81 \quad KGE = 1 - \sqrt{(r-1)^2 + (\alpha-1)^2 + (\beta-1)^2} \quad (S7)$$

82 where r is the correlation coefficient between simulated and observed streamflow; α is the ratio
 83 of their standard deviations; and β is the ratio of their means. KGE ranges from negative infinity
 84 to 1 with values closer to 1 indicating better performance.

85 3) Percent continuous rank probability score reduction (PSR)

86 Continuous rank probability score (CRPS) measures the deviation of the cumulative
 87 distribution function (CDF) of an ensemble from that of a reference (observation in the real-data
 88 case or truth in the synthetic case) [Hersbach, 2000]. If we assume the reference has zero
 89 uncertainty, then its CDF (denoted by $F_t^r(s)$ where t denotes timestep and r denotes observation)
 90 is a unit step function:

$$91 \quad F_t^o(s) = \begin{cases} 0, & s < y_t^r \\ 1, & s \geq y_t^r \end{cases} \quad (S8)$$

92 where y_t^r is the reference variable value and s is a random variable. CRPS is then calculated as
 93 the temporal mean of the CDF deviation from the reference:

$$94 \quad CRPS = \frac{1}{n} \sum_{t=1}^n \int_{-\infty}^{\infty} [F_t^a(s) - F_t^r(s)]^2 ds \quad (S9)$$

95 where $F_t^a(s)$ denotes the CDF of an analysis ensemble at time t , and n is the total length of the
 96 time series. In practice, the continuous CDF of an analysis, $F_t^a(s)$, is empirically estimated by
 97 the finite ensemble. Note that CRPS penalizes both a deviation of the ensemble mean from the
 98 observation and a large ensemble spread. A smaller (i.e., closer-to-zero) CRPS value indicates
 99 better ensemble performance.

100 Analogous to PER, the percent CRPS reduction (PSR) quantifies the percent reduction in
 101 CRPS of a DA analysis compared to the open-loop baseline (note that here the baseline is the
 102 open-loop ensemble instead of the deterministic open-loop run):

$$103 \quad PSR = \left[1 - \frac{CRPS_a}{CRPS_{open}} \right] \times 100 \quad (S10)$$

104 4) Normalized ensemble skill (NENSK)

105 NENSK measures the ensemble-mean error normalized by ensemble spread:

$$106 \quad NENSK = \frac{ENSK}{ENSP} \quad (S11)$$

107 where the ensemble skill (ENSK) and ensemble spread (ENSP) are calculated as:

$$108 \quad ENSK = \frac{1}{n} \sum_{t=1}^n (\bar{y}_t^a - y_t^o)^2 \quad (S12)$$

109
$$ENSP = \frac{1}{n} \sum_{i=1}^n (y_t^{a,(i)} - \bar{y}_t^a)^2 \quad (S13)$$

110 where \bar{y}_t^a denotes the ensemble-mean value at timestep t , and $y_t^{a,(j)}$ denotes the value of the j th
111 ensemble member. In an ideal situation where an ensemble is a correct representation of analysis
112 uncertainty, the observed or true condition looks like one of the realizations of the ensemble
113 [Anderson, 1996; Wilks, 2011] and therefore NENSK should be approximately one [Talagrand
114 et al., 1997; Wilks, 2011] (NENSK > 1 indicates an under-dispersed ensemble while NENSK < 1
115 indicates an over-dispersed ensemble). This metric has been used to verify ensemble hydrologic
116 variables [e.g., De Lannoy et al., 2006; Brocca et al., 2012; Alvarez-Garreton et al., 2014].

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118 **References:**

- 119 Alvarez-Garreton, C., D. Ryu, A. W. Western, W. T. Crow, and D. E. Robertson (2014), The
120 impacts of assimilating satellite soil moisture into a rainfall-runoff model in a semi-arid
121 catchment, *J. Hydrol.*, 519, 2763-2774, doi:10.1016/j.jhydrol.2014.07.041.
- 122 Anderson, J. L. (1996), A method for producing and evaluating probabilistic forecasts from
123 ensemble model integrations, *J. Clim.*, 9, 1518-1530, doi: 10.1175/1520-
124 0442(1996)009<1518:AMFPAE>2.0.CO;2.
- 125 Brocca, L., T. Moramarco, F. Melone, W. Wagner, S. Hasenauer, and S. Hahn (2012),
126 Assimilation of surface-and root-zone ASCAT soil moisture products into rainfall-runoff
127 modeling, *IEEE Trans. Geosci. Remote Sens.*, 50(7), 2542-2555, doi:
128 10.1109/TGRS.2011.2177468.

129 Chen, F., W. T. Crow, and D. Ryu (2014), Dual forcing and state correction via soil moisture
130 assimilation for improved rainfall–runoff modeling, *J. Hydrometeorol.*, 15(5), 1832–
131 1848, doi:10.1175/JHM-D-14-0002.1.

132 Crow, W. T., and D. Ryu (2009), A new data assimilation approach for improving hydrologic
133 prediction using remotely-sensed soil moisture retrievals, *Hydrol. Earth Syst. Sci.*, 12(1-
134 16), doi:10.5194/hess-13-1-2009.

135 De Lannoy, G. J. M., P. R. Houser, V. R. N. Pauwels, and N. E. C. Verhoest (2006), Assessment
136 of model uncertainty for soil moisture through ensemble verification, *J. Geophys.*
137 *Res.*, 111, D10101, doi: 10.1029/2005JD006367.

138 Evensen, G. (1994), Sequential data assimilation with a nonlinear quasi-geostrophic model using
139 Monte Carlo methods to forecast error statistics, *J. Geophys. Res.*, 99(C5), 10143-10162,
140 doi:10.1029/94JC00572.

141 Gupta, H. V., H. Kling, K. K. Yilmaz, and G. F. Martinez (2009), Decomposition of the mean
142 squared error and NSE performance criteria: Implications for improving hydrological
143 modelling, *J. Hydrol.*, 377, 80-91, doi: 10.1016/j.jhydrol.2009.08.003.

144 Hersbach, H. (2000), Decomposition of the continuous ranked probability score for ensemble
145 prediction systems, *Weather Forecast.*, 15(5), 559-570, doi: 10.1175/1520-
146 0434(2000)015<0559:DOTCRP>2.0.CO;2.

147 Massari, C., L. Brocca, A. Tarpanelli, and T. Moramarco (2015), Data Assimilation of Satellite
148 Soil Moisture into Rainfall-Runoff Modelling: A Complex Recipe?, *Remote Sens.*, 7,
149 11403-11433, doi:10.3390/rs70911403.

150 Talagrand, O., R. Vautard, and B. Strauss (1997), Evaluation of probabilistic prediction systems,
151 technical report, Eur. Cent. for Medium-Range Weather Forecast., Reading, UK.

- 152 Wilks, D. S. (2011), Statistical methods in the atmospheric sciences (3rd edition),
153 Elsevier/Academic Press, Amsterdam; Boston.